



JYVÄSKYLÄN YLIOPISTO  
UNIVERSITY OF JYVÄSKYLÄ

## The geometry of anisotropy

Math + X symposium  
on inverse problems and deep learning,  
mitigating natural hazards

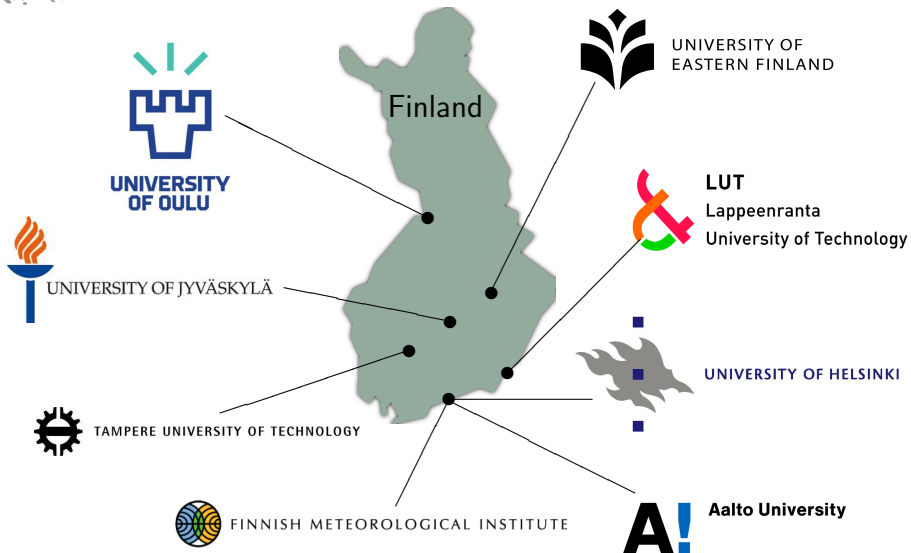
**Joonas Ilmavirta**

January 31, 2020

Based on joint work with

Maarten V. de Hoop, Einar Iversen, Matti Lassas, Teemu Saksala, Bjørn Ursin

# Finnish Centre of Excellence in Inverse Modelling and Imaging 2018-2025



# Introduction

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  - Describing a geometric way to view anisotropy.

- 1 Gravitation
  - Newton's theory
  - Einstein's theory
  - The goal
- 2 Elastic geometry
- 3 Examples

# Newton's theory

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- The gravitational force exerted by the Sun causes the Earth's trajectory to curve.
- The force is described by a simple formula and the equation of motion is an ODE in  $\mathbb{R}^n$ .
- The Newtonian approach is straightforward to use and often a good model.

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- There is a relatively simple equation of motion for the planet: The geodesic equation is a non-linear ODE.
- There is a complicated equation of motion for the geometry itself: Einstein's field equation is a non-linear system of coupled PDEs.
- This model is harder to use but can reach phenomena inaccessible to Newtonian gravity and provides a more geometric way to see the essential structures.



# The goal

A geometric theory of elasticity?

- 1 Gravitation
- 2 Elastic geometry
  - Distance
  - Ray tracing
  - Anisotropy
  - Manifolds to model anisotropy
  - Inverse problems
- 3 Examples

# Distance

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- Distance is measured in units of time.

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  - Traditional view: The trajectory of the particle is curved because wave speed varies.
  - Newer view: The particle goes straight in a curved geometry (geodesic), and the geometry is curved by variations in wave speed.

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- In elastic geometry we measure distance in travel time, and the waves go straight in this geometry.
- Fermat’s principle: The “particles” corresponding to elastic waves go straight in the geometry given by travel time.
- Fermat’s principle is about going straight in the relevant geometry, not about taking the shortest path. These are not the same thing over long distances or shear waves.

# Anisotropy

- Let us ignore polarizations for now.

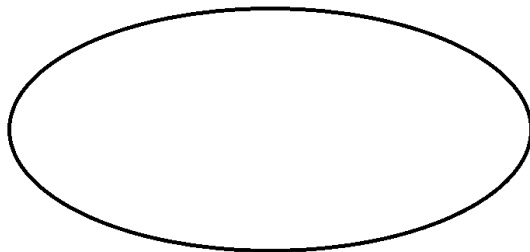


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SPHERE

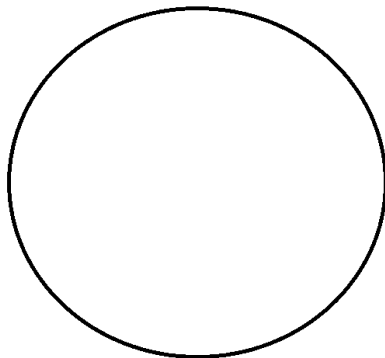


FASTER  
IN THIS  
DIRECTION

SLOWER HERE

Anisotropy.

ISOTROPIC SPHERE



YOU CAN GO EQUALLY FAR  
IN A GIVEN AMOUNT OF TIME,  
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- Sometimes it is more convenient to look at phase velocity.
- The cosphere (the slowness surface) describes the reciprocal of phase velocity.

# Anisotropy

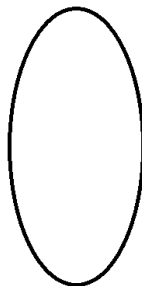
SPHERE



HIGH  
SPEED

LOW  
SLOWNESS

COSPHERE



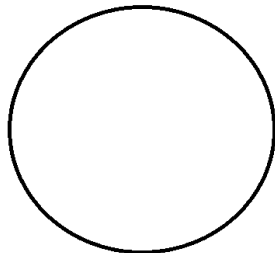
Sphere and cosphere, anisotropic.



SPHERE



COSPHERE



BOTH SAME  
IN ALL  
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Sphere and cosphere, isotropic.

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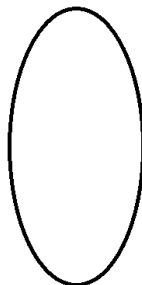
- Isotropy = the sphere and cosphere are spheres.
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- Elliptic anisotropy = the sphere and cosphere are ellipsoids.

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SPHERE



COSPHERE



HIGH  
SPEED

LOW  
SLOWNESS

(BOTH ELLIPSES/ELLIPSOIDS)

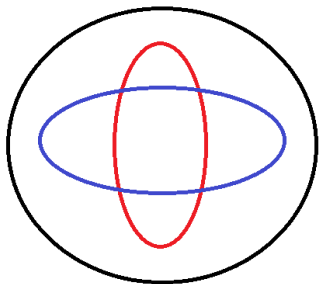
Sphere and cosphere, elliptically anisotropic.

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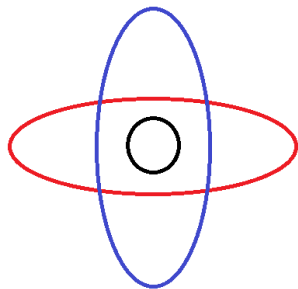
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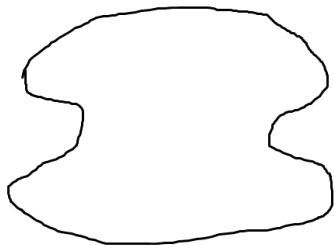
Three polarizations, all elliptically anisotropic.



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- If the cosphere is not convex, the sphere can branch.

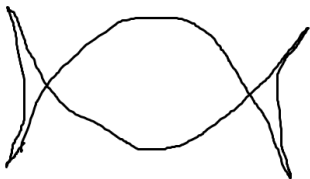
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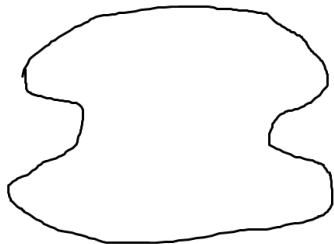
A non-convex cosphere.

# Anisotropy

SPHERE



COSPHERE



A branched sphere.

# Manifolds to model anisotropy

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Multiple metric structures on the same manifold: Each polarization has its own geometry and there is the Euclidean spatial geometry.

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- Elastic Finsler geometry has a decent balance between tractability and applicability.

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- From the slowness surface one can then find the material parameters — the components of the stiffness tensor.

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- From the cosphere bundle one can tell whether the material is isotropic, elliptically anisotropic, completely anisotropic, or even fails to correspond to a stiffness tensor.

1 Gravitation

2 Elastic geometry

3 Examples

- Distance function (de Hoop, Lassas, Saksala)
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- If  $F$  is fiberwise real analytic (elasticity or Riemann!), then  $F$  is determined uniquely.

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- Global uniqueness is can be done with added assumptions: reversibility (point symmetry) and foliation.
- Almost no assumptions are needed in the Riemannian case (Kurylev–Lassas–Uhlmann, 2010).

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- Variations in position ( $Q$ ) and momentum ( $P$ ) satisfy an equation

$$\partial_t \begin{pmatrix} Q(t) \\ P(t) \end{pmatrix} = \begin{pmatrix} W^T(t) & V(t) \\ -U(t) & -W(t) \end{pmatrix} \begin{pmatrix} Q(t) \\ P(t) \end{pmatrix}.$$

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  - Ray-centered coordinates which are more complicated to use but more structure arises.
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- Variations in position ( $Q$ ) and momentum ( $P$ ) satisfy an equation

$$\partial_t \begin{pmatrix} Q(t) \\ P(t) \end{pmatrix} = \begin{pmatrix} W^T(t) & V(t) \\ -U(t) & -W(t) \end{pmatrix} \begin{pmatrix} Q(t) \\ P(t) \end{pmatrix}.$$

- Written in terms of a Jacobi field  $J$  and its covariant derivative, we have instead

$$D_t \begin{pmatrix} J(t) \\ D_t J(t) \end{pmatrix} = \begin{pmatrix} 0 & I \\ -R(t) & 0 \end{pmatrix} \begin{pmatrix} J(t) \\ D_t J(t) \end{pmatrix}.$$

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This approach can hopefully give you:

- A new way to think about anisotropy.
- A new way to encode anisotropy in modeling and computation.

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## 4 The elastic wave equation

- The stiffness tensor
- The elastic wave equation
- The principal symbol
- Polarization
- Singularities and the slowness surface
- Elastic Finsler manifolds



# The stiffness tensor

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- The tensor is very symmetric ( $c_{ijkl} = c_{jikl} = c_{ijlk} = c_{klij}$ ) and quite positive ( $c_{ijkl}\alpha_i\beta_j\beta_k\alpha_l \gtrsim |\alpha|^2|\beta|^2$ ).

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- We will also encounter the density normalized stiffness tensor  $a_{ijkl}(x) = c_{ijkl}(x)/\rho(x)$ .

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- Using Newton's second law with a restoring force given by Hooke's law leads to the elastic wave equation (EWE)

$$\partial_j [c_{ijkl}(x) \partial_k u_l(x, t)] - \rho(x) \partial_t^2 u_i(x, t) = 0,$$

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- If the material is anisotropic ( $c$  is no more symmetric than necessary), then the vector nature of the equation cannot be ignored.
- Elastic waves arising from earthquakes satisfy this equation away from the focus of the event to great accuracy.

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- The principal symbol of the EWE is  $\Gamma(x, \xi) - \omega^2 I$ , where  $\xi = \omega p$ .



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- Polarization vectors are eigenvectors of the Christoffel matrix  $\Gamma$ , so they are orthogonal.
- Decomposition to polarizations only works on the level of singularities. The individual polarizations do not satisfy PDEs.

# Singularities and the slowness surface

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- The admissible slowness vectors are on the slowness surface given by the equation

$$\det(\Gamma(x, p) - I) = 0.$$

# Elastic Finsler manifolds

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- Let  $\lambda(x, p)$  be the largest eigenvalue of  $\Gamma(x, p)$ . The largest eigenvalue corresponds to fastest singularity (qP).
- The qP singularities follow the Hamiltonian flow of  $\lambda: T^*M \rightarrow \mathbb{R}$ .

- The function  $\lambda(x, \cdot): T_x^*\mathbb{R}^3 \rightarrow [0, \infty)$  is smooth outside the origin, strictly convex, and 2-homogeneous.



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- Slowness is a covector and the corresponding vector is the group velocity.

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- Declaring travel time as distance would have defined the same geometry, but in a more implicit manner.

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