

Towards a mathematical theory of seismic tomography on Mars

IAS workshop at HKUST Inverse Problems, Imaging and Partial Differential Equations

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Based on joint work with Maarten de Hoop and Vitaly Katsnelson





The annual Finnish inverse problems conference "Inverese Days" will be organized in Jyväskylä this year.

16-18 December, 2019.

https://www.jyu.fi/science/en/maths/research/ inverse-problems/id2019/inverse-days-2019

http://r.jyu.fi/yVK

All kinds of inverse problems in all fields are welcome!

Goals

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- Assume perfect measurements from a single ideal seismometer. What can you say for sure and is there an inversion algorithm?
- What would be useful data sets for future missions?
- Grand goal: A mathematical theory of seismic planetary exploration.

Seeing the radial Martian mantle with InSight

Seeing the entire planet

A small but reliable step

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- Mars is roughly spherically symmetric. There are reliable ways to reconstruct a radial model of the (upper) mantle from a single station. (The mantle determines the CMB.)
- I will ignore noise, model errors, finiteness, stability, and many other practical things.

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- Data: Pairs of directions (\approx angle from normal) and times. Uknown: Wave speed (\approx geometry).
- The set of all periodic travel times is the length spectrum.

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- Reconstructing the wave speed from travel time data is hard, even with data everywhere on the surface.
- Solution: Linearize!
- Linearized data: Pairs of periodic broken rays and integrals over them. Uknown: Variations of wave speed (a function).



Periodic seismic ray reflecting on the surface and CMB.

Joonas Ilmavirta (University of Jyväskylä)

Seismology on Mars

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Solving the linearized problem gives an iterative algorithm to solve the nonlinear one. (Uniqueness should be provable for the non-linear one, too.)

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- The oscillations are excited by marsquakes, atmosphere, meteorite impacts, and other possible events.
- The oscillations can be decomposed into eigenmodes which have their own frequencies.
- The different modes are excited differently in different events, but one thing remains: the set of frequencies — the spectrum of free oscillations. (We are at first interested in properties of the planet, not properties of the events.)
- The spectrum of free oscillations can be measured from any single point.
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- Again wave speed = geometry!

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Question

If a family of wave speeds $c_s(r)$ have the same spectrum, are the equal? Is the (Martian) mantle spectrally rigid?

Consider the annulus (mantle) $M = \overline{B}(0,1) \setminus B(0,R) \subset \mathbb{R}^3$. Let $c_s(r)$ be a family of radial sound speeds depending C^{∞} -smoothly on both $s \in (-\varepsilon, \varepsilon)$ and $r \in [R, 1]$. Assume each c_s satisfies the Herglotz condition and a generic geometrical condition.

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This simple model of the round Martian mantle is spectrally rigid!

Let $\lambda_0 < \lambda_1 \leq \lambda_2 \leq \dots$ be the positive eigenvalues of the Laplace–Beltrami operator. Define a function $f : \mathbb{R} \to \mathbb{R}$ by

$$f(t) = \sum_{k=0}^{\infty} \cos\left(\sqrt{\lambda_k} \cdot t\right).$$

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In particular, the spectrum determines the length spectrum. It suffices to prove length spectral rigidity.

Method B: Spectral data



The length spectrum is \mathbb{Z} .

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Method B: Spectral data



Trace function $f(t) = \sum_k \cos(\sqrt{\lambda_k} \cdot t)$ computed from k = 0, 1, 2, 3, 4.

Method B: Spectral data



The trace computed from the spectrum of free oscillations in PREM. Singularities are visible.

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The data set is independent although the method is related.

Method C: Meteorite impacts

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- Multiple arrivals or a priori information tells the time *T* around the great circle.

Method C: Meteorite impacts



Two surface wave arrivals from the same event.

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Seismology on Mars

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- Assuming the seismometer can detect directions of surface wave arrivals, we can deduce the time and place of the event.
- This was all done on surface, and it gives rise to interior data: Now using body waves we know the travel time between InSight and the source.
- To get here, we needed to assume spherical symmetry only on the surface, but the arising problem is easiest to solve if the symmetry extends inside.
Method C: Meteorite impacts



The body wave whose initial point and time were located with surface waves.

Joonas Ilmavirta (University of Jyväskylä)

Seismology on Mars

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- The linearized problem is X-ray tomography (or an Abel transform), and can also be solved explicitly. (e.g. de Hoop–I., 2017)





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- In the Earth the Herglotz condition is satisfied in the whole mantle for both P and S. On Mars it will at least hold in the upper mantle.
- The three methods use independently obtained datasets.
- If the three reconstructions all work and give similar results, we can be quite confident.
- This gives us an isotropic radially symmetric reference model of the mantle, which is a stepping stone towards deeper and finer structure.

Summary



Three ways to see the mantle from InSight.

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Seismology on Mars

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- A: From noise correlations to (linearized) travel times.
- B: From spectrum to length spectrum.
- C: Meteorites; body wave data calibrated by surface waves.



Seeing the radial Martian mantle with InSight

Seeing the entire planet

Spectral perturbation theory

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- Proving precise results outside spherical symmetry with one measurement point is hard.
- A natural approach to small lateral inhomogeneities is perturbation theory with respect to to a spherically symmetric reference model.
- In a simple (scalar) model, the medium is described by a single wave speed c(x) and the spectrum depends on it: Sp(c).
- We write the wave speed as a function of a parameter, $c_s(x)$, and expand the spectrum in *s*:

$$Sp(c_s) = Sp(c_0) + sL(\delta c) + \mathcal{O}(s^2),$$

where $\delta c = \frac{d}{ds}c_s|_{s=0}$, *L* is the Gâteaux derivative of the spectrum, and '+' is roughly a plus.

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- If the reference medium c_0 is known and s is small, it is sufficient(ish) to invert the linear operator L.
- On Mars, c_0 would be the radial reference model.
- The better the radial (or other initial) guess is, the better the perturbation theory works.
- The perturbation δc can be expanded in spherical harmonics and the operator *L* can be written fairly explicitly.

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First observations:

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It is best to start with a scalar model in 2D, not a fully polarized 3D model.

Joonas Ilmavirta (University of Jyväskylä)

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- We assumed that the surface is spherically symmetric (or otherwise known), but we needed no assumption on the interior.
- This leads to travel time data: The travel times (geometrically: distances) are known from all points on the surface to a single fixed point.

Half-local X-ray tomography



The body wave whose initial point and time were located with surface waves.

Joonas Ilmavirta (University of Jyväskylä)

Seismology on Mars

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Question

What if the point is replaced by a small open set — a detector array?

Half-local X-ray tomography



Boundary distance rigidity: Do the distances between all boundary points determine the geometry?

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Seismology on Mars

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Half-local X-ray tomography



We have an accessible region — a measurement array. The size is exaggerated.

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Seismology on Mars

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Half-local X-ray tomography



In the local boundary distance problem one knows the distances between the points in the small set and wants to find the geometry near that set.

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Half-local X-ray tomography



The "half-local" boundary distance data has more information and one wants to reconstruct the whole geometry.

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Seismology on Mars

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This is possible in Euclidean geometry or with real analytic perturbations but always unstable.

Sources and receivers

Joonas Ilmavirta (University of Jyväskylä)

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- What is the minimal number of measurement points for uniqueness?



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- Most geometrical inverse problems work with smooth manifolds. How to add conormal singularities and finite interior regularity?
- How does spectral rigidity and X-ray tomography work in an onion?

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- ... and even Finsler is not enough for all polarizations.

Geometry of periodic geodesic

Put any metric on the unit sphere and fix a point on it. How many directions are there so that the geodesic will make it back to the point?

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How does X-ray tomography change when Anosov flow is replaced by dispersing billiards?

A theory?

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