

Dix's inverse problem on elastic Finsler manifolds

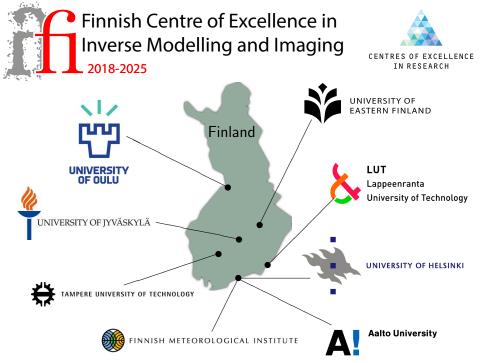
Geo-Mathematical Imaging Group 2019 project review

Joonas Ilmavirta

April 29, 2019

Based on joint work with Maarten de Hoop and Matti Lassas





Goals

• Understanding how and why to model elastic waves in terms of Finsler geometry.

- Understanding how and why to model elastic waves in terms of Finsler geometry.
- Seeing how geometrical tools can be used to solve Dix's problem in great generality.



Dix's inverse problem

Joonas Ilmavirta (University of Jyväskylä)

In linear elasticity the displacement vector $\boldsymbol{u}(\boldsymbol{x},t)$ satisfies the elastic wave equation

$$\partial_j [c_{ijkl}(x)\partial_k u_l(x,t)] - \rho(x)\partial_t^2 u_i(x,t) = 0,$$

where $c_{ijkl}(x)$ is the stiffness tensor.

In linear elasticity the displacement vector u(x,t) satisfies the elastic wave equation

$$\partial_j [c_{ijkl}(x)\partial_k u_l(x,t)] - \rho(x)\partial_t^2 u_i(x,t) = 0,$$

where $c_{ijkl}(x)$ is the stiffness tensor.

This is a good model for low amplitude body waves in the Earth.

In linear elasticity the displacement vector u(x,t) satisfies the elastic wave equation

$$\partial_j [c_{ijkl}(x)\partial_k u_l(x,t)] - \rho(x)\partial_t^2 u_i(x,t) = 0,$$

where $c_{ijkl}(x)$ is the stiffness tensor.

This is a good model for low amplitude body waves in the Earth.

Instead of looking at all of a solution u(x,t), we look at singularities (non-smooth component) of the solutions. This corresponds roughly to the high frequency limit.

In linear elasticity the displacement vector u(x,t) satisfies the elastic wave equation

$$\partial_j [c_{ijkl}(x)\partial_k u_l(x,t)] - \rho(x)\partial_t^2 u_i(x,t) = 0,$$

where $c_{ijkl}(x)$ is the stiffness tensor.

This is a good model for low amplitude body waves in the Earth.

Instead of looking at all of a solution u(x,t), we look at singularities (non-smooth component) of the solutions. This corresponds roughly to the high frequency limit.

The singularities can be described in geometrical terms as point particles or wave fronts.

Propagation of singularities

We can think of singularities as point particles with a polarization vector attached to them. (Microlocal analysis makes this idea precise.)

We can think of singularities as point particles with a polarization vector attached to them. (Microlocal analysis makes this idea precise.)

The slowness vector p (inverse phase velocity), position x, and polarization A of the singularity have to satisfy

$$[\Gamma(x,p) - I]A = 0,$$

where $\Gamma_{ij}(x,p) = \rho^{-1} c_{iklj} p_k p_l$ is the Christoffel matrix.

We can think of singularities as point particles with a polarization vector attached to them. (Microlocal analysis makes this idea precise.)

The slowness vector p (inverse phase velocity), position x, and polarization A of the singularity have to satisfy

$$[\Gamma(x,p) - I]A = 0,$$

where $\Gamma_{ij}(x,p) = \rho^{-1} c_{iklj} p_k p_l$ is the Christoffel matrix.

The possible values of \boldsymbol{p} are on the slowness surface defined by the equation

$$\det[\Gamma(x,p) - I] = 0.$$

The slowness surface describes what the singularities can look like, but not how they move.

- The slowness surface describes what the singularities can look like, but not how they move.
- Microlocal analysis (or physical reasoning) tells how singularities propagate: They follow a Hamiltonian flow related to the slowness surface.

- The slowness surface describes what the singularities can look like, but not how they move.
- Microlocal analysis (or physical reasoning) tells how singularities propagate: They follow a Hamiltonian flow related to the slowness surface.
- If the slowness surface is smooth and strictly convex (qP!), then this coincides with the geodesic flow of a Finsler manifold, where the Finsler metric is the Legendre transform of the slowness surface.

Propagation of singularities

Our view: Elastic waves move straight in a curved space whose geometry is determined by the sound speed.

Our view: Elastic waves move straight in a curved space whose geometry is determined by the sound speed. (Interaction is encoded in geometry.)

Our view: Elastic waves move straight in a curved space whose geometry is determined by the sound speed. (Interaction is encoded in geometry.)

Distance is the basis of any geometry.

Our view: Elastic waves move straight in a curved space whose geometry is determined by the sound speed. (Interaction is encoded in geometry.)

Distance is the basis of any geometry. In this elastic geometry distance is measured in travel time.

Our view: Elastic waves move straight in a curved space whose geometry is determined by the sound speed. (Interaction is encoded in geometry.)

Distance is the basis of any geometry. In this elastic geometry distance is measured in travel time.

If the slowness surfaces are ellipsoids, we have a Riemannian manifold.

Our view: Elastic waves move straight in a curved space whose geometry is determined by the sound speed. (Interaction is encoded in geometry.)

Distance is the basis of any geometry. In this elastic geometry distance is measured in travel time.

If the slowness surfaces are ellipsoids, we have a Riemannian manifold. But we do not assume any kind of isotropy!

From elasticity to Finsler geometry

• Elasticity is described in terms of a stiffness tensor.

- Elasticity is described in terms of a stiffness tensor.
- The stiffness tensor determines a Christoffel matrix, which in turn determines the slowness surface.

- Elasticity is described in terms of a stiffness tensor.
- The stiffness tensor determines a Christoffel matrix, which in turn determines the slowness surface.
- The slowness surface corresponds to a co-Finsler metric.

- Elasticity is described in terms of a stiffness tensor.
- The stiffness tensor determines a Christoffel matrix, which in turn determines the slowness surface.
- The slowness surface corresponds to a co-Finsler metric.
- Taking the Legendre transform gives a Finsler metric which corresponds to group velocities.

- Elasticity is described in terms of a stiffness tensor.
- The stiffness tensor determines a Christoffel matrix, which in turn determines the slowness surface.
- The slowness surface corresponds to a co-Finsler metric.
- Taking the Legendre transform gives a Finsler metric which corresponds to group velocities.
- The singularities of elastic waves (= elastic waves as point particles) follow geodesics of this Finsler geometry.

Inverse problems

Problem (Physical inverse problem)

Given some boundary measurements, find the elastic parameters.

Problem (Physical inverse problem)

Given some boundary measurements, find the elastic parameters.

Problem (Mathematical inverse problem)

Given some boundary measurements, find the elastic geometry.

Problem (Physical inverse problem)

Given some boundary measurements, find the elastic parameters.

Problem (Mathematical inverse problem)

Given some boundary measurements, find the elastic geometry.

Instead of focusing on the reduced stiffness tensor $a_{ijkl} = \rho^{-1}c_{ijkl}$, we try to find the slowness surface at every point.

Problem (Physical inverse problem)

Given some boundary measurements, find the elastic parameters.

Problem (Mathematical inverse problem)

Given some boundary measurements, find the elastic geometry.

Instead of focusing on the reduced stiffness tensor $a_{ijkl} = \rho^{-1}c_{ijkl}$, we try to find the slowness surface at every point. Getting from the slowness surface back to the stiffness tensor is not trivial.

Problem (Physical inverse problem)

Given some boundary measurements, find the elastic parameters.

Problem (Mathematical inverse problem)

Given some boundary measurements, find the elastic geometry.

Instead of focusing on the reduced stiffness tensor $a_{ijkl} = \rho^{-1}c_{ijkl}$, we try to find the slowness surface at every point. Getting from the slowness surface back to the stiffness tensor is not trivial.

Gains:

Freedom to have any shape of slowness surfaces, not necessarily arising from Hookean elasticity.

Problem (Physical inverse problem)

Given some boundary measurements, find the elastic parameters.

Problem (Mathematical inverse problem)

Given some boundary measurements, find the elastic geometry.

Instead of focusing on the reduced stiffness tensor $a_{ijkl} = \rho^{-1}c_{ijkl}$, we try to find the slowness surface at every point. Getting from the slowness surface back to the stiffness tensor is not trivial.

Gains:

- Freedom to have any shape of slowness surfaces, not necessarily arising from Hookean elasticity.
- Access to powerful tools of differential geometry.

Changes of coordinates

Joonas Ilmavirta (University of Jyväskylä)

Consider some domain $\Omega \subset \mathbb{R}^3$ (e.g. the Earth) and a Finsler metric F on it. (It is a function $T\Omega \to \mathbb{R}$.)

Consider some domain $\Omega \subset \mathbb{R}^3$ (e.g. the Earth) and a Finsler metric F on it. (It is a function $T\Omega \to \mathbb{R}$.)

Take any diffeomorphism (change of coordinates) $\phi \colon \overline{\Omega} \to \overline{\Omega}$ so that $\phi(x) = x$ for all $x \in \partial \Omega$. This can be any smooth distortion of the planet.

Consider some domain $\Omega \subset \mathbb{R}^3$ (e.g. the Earth) and a Finsler metric F on it. (It is a function $T\Omega \to \mathbb{R}$.)

Take any diffeomorphism (change of coordinates) $\phi \colon \overline{\Omega} \to \overline{\Omega}$ so that $\phi(x) = x$ for all $x \in \partial \Omega$. This can be any smooth distortion of the planet.

One can define (and write) a new Finsler metric $\phi^*F,$ the pullback of F over $\phi.$

Consider some domain $\Omega \subset \mathbb{R}^3$ (e.g. the Earth) and a Finsler metric F on it. (It is a function $T\Omega \to \mathbb{R}$.)

Take any diffeomorphism (change of coordinates) $\phi : \overline{\Omega} \to \overline{\Omega}$ so that $\phi(x) = x$ for all $x \in \partial \Omega$. This can be any smooth distortion of the planet.

One can define (and write) a new Finsler metric ϕ^*F , the pullback of F over ϕ . All geodesics and everything behaves exactly the same — F and ϕ^*F look exactly alike from the surface.

Changes of coordinates

The freedom to change coordinates is helpful in geometry but problematic in geophysics: It seems that even if one can find the elastic structure, one cannot tell whether the coordinates are Euclidean.

The best one can hope for is uniqueness of F up to changes of coordinates.

The best one can hope for is uniqueness of F up to changes of coordinates.

This is known to happen in anisotropic electrical impedance tomography and anisotropic electromagnetism: the geometrical gauge freedom of diffeomorphisms is inevitable.

The best one can hope for is uniqueness of F up to changes of coordinates.

This is known to happen in anisotropic electrical impedance tomography and anisotropic electromagnetism: the geometrical gauge freedom of diffeomorphisms is inevitable.

But we do not know whether this freedom is there in elasticity. There are reasons to believe that elasticity would be more sensitive to Euclidean geometry than electromagnetism, but we have no proof — yet.

The best one can hope for is uniqueness of F up to changes of coordinates.

This is known to happen in anisotropic electrical impedance tomography and anisotropic electromagnetism: the geometrical gauge freedom of diffeomorphisms is inevitable.

But we do not know whether this freedom is there in elasticity. There are reasons to believe that elasticity would be more sensitive to Euclidean geometry than electromagnetism, but we have no proof — yet.

Not all Finsler metrics arise from a stiffness tensor. When does ϕ^*F do?





Dix's inverse problem

Problem (Dix's geophysical inverse problem)

If you measure the wave fronts coming from interior (real or virtual) point sources, can you find the properties of the medium?

Problem (Dix's geophysical inverse problem)

If you measure the wave fronts coming from interior (real or virtual) point sources, can you find the properties of the medium?

Problem (Dix's geometrical inverse problem)

If you measure the wave fronts of the geodesic flow coming from points in a Finsler manifold, can you find the manifold and its Finsler metric?

Problem (Dix's geophysical inverse problem)

If you measure the wave fronts coming from interior (real or virtual) point sources, can you find the properties of the medium?

Problem (Dix's geometrical inverse problem)

If you measure the wave fronts of the geodesic flow coming from points in a Finsler manifold, can you find the manifold and its Finsler metric?

There is an unknown Finsler manifold (M, F) and measurements are conducted in an open set $U \subset M$. The set U can be tiny.

Problem (Dix's geophysical inverse problem)

If you measure the wave fronts coming from interior (real or virtual) point sources, can you find the properties of the medium?

Problem (Dix's geometrical inverse problem)

If you measure the wave fronts of the geodesic flow coming from points in a Finsler manifold, can you find the manifold and its Finsler metric?

There is an unknown Finsler manifold (M, F) and measurements are conducted in an open set $U \subset M$. The set U can be tiny.

The Riemannian version was solved by de Hoop, Holman, Iversen, Lassas, and Ursin (2015).

Problem (Dix's geophysical inverse problem)

If you measure the wave fronts coming from interior (real or virtual) point sources, can you find the properties of the medium?

Problem (Dix's geometrical inverse problem)

If you measure the wave fronts of the geodesic flow coming from points in a Finsler manifold, can you find the manifold and its Finsler metric?

There is an unknown Finsler manifold (M, F) and measurements are conducted in an open set $U \subset M$. The set U can be tiny.

The Riemannian version was solved by de Hoop, Holman, Iversen, Lassas, and Ursin (2015). Moving to Finsler geometry means that we can deal with arbitrary (non-elliptic) anisotropy.

Problem (Dix's geophysical inverse problem)

If you measure the wave fronts coming from interior (real or virtual) point sources, can you find the properties of the medium?

Problem (Dix's geometrical inverse problem)

If you measure the wave fronts of the geodesic flow coming from points in a Finsler manifold, can you find the manifold and its Finsler metric?

There is an unknown Finsler manifold (M, F) and measurements are conducted in an open set $U \subset M$. The set U can be tiny.

The Riemannian version was solved by de Hoop, Holman, Iversen, Lassas, and Ursin (2015). Moving to Finsler geometry means that we can deal with arbitrary (non-elliptic) anisotropy.

We make no assumptions on symmetry, isotropy, or homogeneity.

Theorem (de Hoop-I.-Lassas)

Let (M, F) be any Finsler manifold and $U \subset M$ an open set. Suppose F is fiberwise analytic. If we know the smooth metric "spheres" (wave fronts from point sources) together with radii (travel times) in the set U, we can find the universal cover of (M, F).

Theorem (de Hoop–I.–Lassas)

Let (M, F) be any Finsler manifold and $U \subset M$ an open set. Suppose F is fiberwise analytic. If we know the smooth metric "spheres" (wave fronts from point sources) together with radii (travel times) in the set U, we can find the universal cover of (M, F).

The geometry (= slowness surfaces at all points) is determined uniquely, but global topology is not.

Theorem (de Hoop–I.–Lassas)

Let (M, F) be any Finsler manifold and $U \subset M$ an open set. Suppose F is fiberwise analytic. If we know the smooth metric "spheres" (wave fronts from point sources) together with radii (travel times) in the set U, we can find the universal cover of (M, F).

The geometry (= slowness surfaces at all points) is determined uniquely, but global topology is not. This is not an issue for the Earth.

Theorem (de Hoop-I.-Lassas)

Let (M, F) be any Finsler manifold and $U \subset M$ an open set. Suppose F is fiberwise analytic. If we know the smooth metric "spheres" (wave fronts from point sources) together with radii (travel times) in the set U, we can find the universal cover of (M, F).

The geometry (= slowness surfaces at all points) is determined uniquely, but global topology is not. This is not an issue for the Earth.

All Finsler metrics arising from elasticity are fiberwise analytic.

Theorem (de Hoop–I.–Lassas)

Let (M, F) be any Finsler manifold and $U \subset M$ an open set. Suppose F is fiberwise analytic. If we know the smooth metric "spheres" (wave fronts from point sources) together with radii (travel times) in the set U, we can find the universal cover of (M, F).

The geometry (= slowness surfaces at all points) is determined uniquely, but global topology is not. This is not an issue for the Earth.

All Finsler metrics arising from elasticity are fiberwise analytic. This is a good example of useful additional structure that comes from good modelling — the model is general enough to handle all anisotropy but narrow enough to exclude nonphysical oddities.

Ideas and methods

Joonas Ilmavirta (University of Jyväskylä)

• Given two triples (*x*, *v*, *t*) in *U*, the data tells whether they hit the same point.

- Given two triples (*x*, *v*, *t*) in *U*, the data tells whether they hit the same point.
- One does not know anything outside *U*, so coordinates have to be built using elements in *U*.

- Given two triples (*x*, *v*, *t*) in *U*, the data tells whether they hit the same point.
- One does not know anything outside *U*, so coordinates have to be built using elements in *U*. Surface normal coordinates are natural.

- Given two triples (*x*, *v*, *t*) in *U*, the data tells whether they hit the same point.
- One does not know anything outside *U*, so coordinates have to be built using elements in *U*. Surface normal coordinates are natural.
- Jacobi fields describe local curvature along a geodesic.

- Given two triples (*x*, *v*, *t*) in *U*, the data tells whether they hit the same point.
- One does not know anything outside *U*, so coordinates have to be built using elements in *U*. Surface normal coordinates are natural.
- Jacobi fields describe local curvature along a geodesic. The data allows us to find all the Jacobi fields along any geodesic through the set *U* by solving a non-linear ODE system.

- Given two triples (*x*, *v*, *t*) in *U*, the data tells whether they hit the same point.
- One does not know anything outside *U*, so coordinates have to be built using elements in *U*. Surface normal coordinates are natural.
- Jacobi fields describe local curvature along a geodesic. The data allows us to find all the Jacobi fields along any geodesic through the set *U* by solving a non-linear ODE system.
- We can find the geometry along any geodesic, and this information must be turned into a global geometrical description.

- Given two triples (*x*, *v*, *t*) in *U*, the data tells whether they hit the same point.
- One does not know anything outside *U*, so coordinates have to be built using elements in *U*. Surface normal coordinates are natural.
- Jacobi fields describe local curvature along a geodesic. The data allows us to find all the Jacobi fields along any geodesic through the set *U* by solving a non-linear ODE system.
- We can find the geometry along any geodesic, and this information must be turned into a global geometrical description. This requires using analyticity and building an atlas (a collection of coordinate charts).

DISCOVERING MATH at JYU. Since 1863.

Slides and papers available at http://users.jyu.fi/~jojapeil