

Spectral rigidity of the round Earth

CCIMI Seminar, University of Cambridge

Joonas Ilmavirta

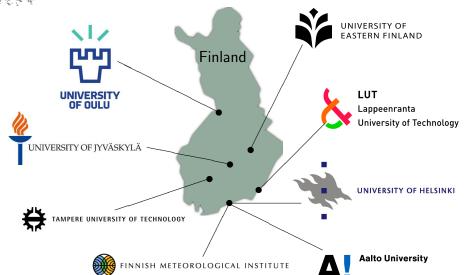
April 9, 2018

Based on joint work with

Maarten de Hoop and Vitaly Katsnelson (Rice)







Conference announcement

A workshop "Inverse problems, PDE and geometry" will be organized in Jyväskylä, Finland, on August 20–23, 2018.

The workshop will focus on recent progress in the mathematical theory of inverse problems and related methods in PDEs, geometry and microlocal analysis.

Scientific committee: Yaroslav Kurylev (UCL), Mikko Salo (Jyväskylä), Gunther Uhlmann (Washington/HKUST)

Can you hear what is inside the Earth?

- Can you hear what is inside the Earth?
- What can one tell about the Earth just by the spectrum of its free oscillations?

- Can you hear what is inside the Earth?
- What can one tell about the Earth just by the spectrum of its free oscillations?
- This is an inverse spectral problem. A hard one.

- Can you hear what is inside the Earth?
- What can one tell about the Earth just by the spectrum of its free oscillations?
- This is an inverse spectral problem. A hard one.
- There is a weaker version of the spectral problem: the spectral rigidity problem.

- Can you hear what is inside the Earth?
- What can one tell about the Earth just by the spectrum of its free oscillations?
- This is an inverse spectral problem. A hard one.
- There is a weaker version of the spectral problem: the spectral rigidity problem.
- Can we solve the simpler problem if we assume the Earth to be spherically symmetric?

- Can you hear what is inside the Earth?
- What can one tell about the Earth just by the spectrum of its free oscillations?
- This is an inverse spectral problem. A hard one.
- There is a weaker version of the spectral problem: the spectral rigidity problem.
- Can we solve the simpler problem if we assume the Earth to be spherically symmetric?
- Yes!

Outline

- Seismic spectral data
 - The spectrum of free oscillations
 - The spectrum of periodic orbits
 - The goal
- Spectra of a manifold with boundary
- 3 Different forms of uniqueness
- Spherical symmetry
- The main results
- 6 Anisotropy and geometry
- Ray transforms in low regularity

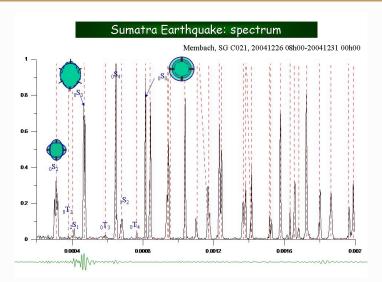
 Like any elastic object, the Earth can vibrate. These vibrations are known as free oscillations.

- Like any elastic object, the Earth can vibrate. These vibrations are known as free oscillations.
- The oscillations are excited (started) by large earthquakes.
 Oscillations are visible once the more violent and transient first stages pass.

- Like any elastic object, the Earth can vibrate. These vibrations are known as free oscillations.
- The oscillations are excited (started) by large earthquakes.
 Oscillations are visible once the more violent and transient first stages pass.
- The amplitudes of different modes vary between different events, but the frequencies are always the same.

- Like any elastic object, the Earth can vibrate. These vibrations are known as free oscillations.
- The oscillations are excited (started) by large earthquakes.
 Oscillations are visible once the more violent and transient first stages pass.
- The amplitudes of different modes vary between different events, but the frequencies are always the same.
- The set of these frequencies (with multiplicity) is the spectrum of free oscillations.

- Like any elastic object, the Earth can vibrate. These vibrations are known as free oscillations.
- The oscillations are excited (started) by large earthquakes.
 Oscillations are visible once the more violent and transient first stages pass.
- The amplitudes of different modes vary between different events, but the frequencies are always the same.
- The set of these frequencies (with multiplicity) is the spectrum of free oscillations.
- About 10 000 first frequencies are known.



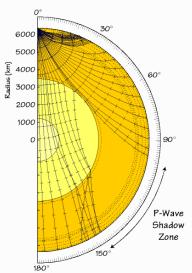
Spectrum of free oscillations from an earthquake.

 One can think of seismic waves in terms of ray theory: Individual points in a seismic wave (front) travel along a certain path.

- One can think of seismic waves in terms of ray theory: Individual points in a seismic wave (front) travel along a certain path.
- Some of the wave paths are periodic. Every periodic wave path has a length (in time).

- One can think of seismic waves in terms of ray theory: Individual points in a seismic wave (front) travel along a certain path.
- Some of the wave paths are periodic. Every periodic wave path has a length (in time).
- The set of all lengths of periodic seismic wave paths is the "length spectrum" of the Earth.

- One can think of seismic waves in terms of ray theory: Individual points in a seismic wave (front) travel along a certain path.
- Some of the wave paths are periodic. Every periodic wave path has a length (in time).
- The set of all lengths of periodic seismic wave paths is the "length spectrum" of the Earth.
- Originally the length spectrum was just a mathematical tool, but it turns out it can be measured directly using deep earthquakes.



Seismic wave paths and the P-wave shadow zone. (Wikimedia Commons)

The goal

Problem

Given the spectrum of free oscillations or the length spectrum of the Earth, reconstruct the Earth.

The goal

Problem

Given the spectrum of free oscillations or the length spectrum of the Earth, reconstruct the Earth.

This problem only makes sense within a given model.

The goal

Problem

Given the spectrum of free oscillations or the length spectrum of the Earth, reconstruct the Earth.

This problem only makes sense within a given model.

We want to reconstruct the Earth in the natural Cartesian coordinates.

Outline

- Seismic spectral data
- Spectra of a manifold with boundary
 - Manifolds with boundary
 - The spectrum of the Laplacian
 - The length spectrum
- Different forms of uniqueness
- Spherical symmetry
- The main results
- 6 Anisotropy and geometry
- Ray transforms in low regularity

• We model the Earth as a Riemannian manifold with boundary.

- We model the Earth as a Riemannian manifold with boundary.
- In practice, the Earth is the closed unit ball $M=\bar{B}(0,1)\subset\mathbb{R}^3$. The anisotropic sound speed is modeled with a Riemannian metric g on M.

- We model the Earth as a Riemannian manifold with boundary.
- In practice, the Earth is the closed unit ball $M=\bar{B}(0,1)\subset\mathbb{R}^3.$ The anisotropic sound speed is modeled with a Riemannian metric g on M.
- Physically, this corresponds to omitting S-waves and including only elliptic anisotropy.

 The modes of free oscillations correspond to Neumann eigenfunctions of the Laplace
—Beltrami operator of (M, g).

- The modes of free oscillations correspond to Neumann eigenfunctions of the Laplace—Beltrami operator of (M, g).
- If the sound speed is isotropic, then $g=c^{-2}e$ and the Laplace–Beltrami operator in dimension n is

$$\Delta_g u(x) = c(x)^n \operatorname{div}(c(x)^{2-n} \nabla u(x)).$$

- The modes of free oscillations correspond to Neumann eigenfunctions of the Laplace—Beltrami operator of (M, g).
- If the sound speed is isotropic, then $g=c^{-2}e$ and the Laplace–Beltrami operator in dimension n is

$$\Delta_g u(x) = c(x)^n \operatorname{div}(c(x)^{2-n} \nabla u(x)).$$

 \bullet The spectrum of free oscillations is the Neumann spectrum of the Laplace–Beltrami operator $\Delta_g.$

A seismic wave path corresponds to a geodesic.

- A seismic wave path corresponds to a geodesic.
- Seismic waves reflect at the surface, so they are in fact billiard trajectories or broken rays.

- A seismic wave path corresponds to a geodesic.
- Seismic waves reflect at the surface, so they are in fact billiard trajectories or broken rays.
- \bullet The length spectrum of (M,g) is the set of all lengths of the periodic broken rays.

Outline

- Seismic spectral data
- Spectra of a manifold with boundary
- Oifferent forms of uniqueness
 - Difficulties
 - Diffeomorphisms and coordinates
 - Global uniqueness
 - Local uniqueness
 - Spectral rigidity
- Spherical symmetry
- The main results
- 6 Anisotropy and geometry
- Ray transforms in low regularity

 The obvious first conjecture is that the spectrum determines the Riemannian manifold (with boundary) uniquely.

- The obvious first conjecture is that the spectrum determines the Riemannian manifold (with boundary) uniquely.
- Proving this conjecture is difficult for two reasons:

- The obvious first conjecture is that the spectrum determines the Riemannian manifold (with boundary) uniquely.
- Proving this conjecture is difficult for two reasons:
 - The required tools do not yet exist on general manifolds with boundary.

- The obvious first conjecture is that the spectrum determines the Riemannian manifold (with boundary) uniquely.
- Proving this conjecture is difficult for two reasons:
 - The required tools do not yet exist on general manifolds with boundary.
 - The conjecture is false.

 The main obstacle to uniqueness is that there are no preferred coordinates.

- The main obstacle to uniqueness is that there are no preferred coordinates.
- If $\phi \colon M \to M$ is a diffeomorphism, then (M,g) and (M,ϕ^*g) give the same spectrum.

- The main obstacle to uniqueness is that there are no preferred coordinates.
- If $\phi \colon M \to M$ is a diffeomorphism, then (M,g) and (M,ϕ^*g) give the same spectrum.
- We can take any change of coordinates whatsoever and use it to distort the metric, but the spectrum stays the same.

- The main obstacle to uniqueness is that there are no preferred coordinates.
- If $\phi \colon M \to M$ is a diffeomorphism, then (M,g) and (M,ϕ^*g) give the same spectrum.
- We can take any change of coordinates whatsoever and use it to distort the metric, but the spectrum stays the same.
- Physically: There are preferred and natural Cartesian coordinates.
 But the anisotropic model is not "sensitive to the underlying Euclidean geometry", so the Cartesian coordinates cannot be recognized. It is impossible to find the metric (elliptically anisotropic sound speed) in Cartesian coordinates from spectral data.

Global uniqueness

Problem

Let g_1 and g_2 be two Riemannian metrics on a manifold M with boundary. If they give the same spectrum, is there a diffeomorphism $\phi \colon M \to M$ so that $g_1 = \phi^* g_2$?

Global uniqueness

Problem

Let g_1 and g_2 be two Riemannian metrics on a manifold M with boundary. If they give the same spectrum, is there a diffeomorphism $\phi \colon M \to M$ so that $g_1 = \phi^* g_2$?

This is too hard.

Local uniqueness

Problem

Let g_1 and g_2 be two Riemannian metrics on a manifold M with boundary. Suppose g_1 is very close to g_2 . If they give the same spectrum, is there a diffeomorphism $\phi \colon M \to M$ so that $g_1 = \phi^* g_2$?

Local uniqueness

Problem

Let g_1 and g_2 be two Riemannian metrics on a manifold M with boundary. Suppose g_1 is very close to g_2 . If they give the same spectrum, is there a diffeomorphism $\phi \colon M \to M$ so that $g_1 = \phi^* g_2$?

This is still too hard.

Problem

Let q_s be family of Riemannian metrics on a manifold M with boundary, depending on a parameter $s \in (-\varepsilon, \varepsilon)$. If they all give the same spectrum, are there a diffeomorphisms $\phi_s \colon M \to M$ so that $g_0 = \phi_s^* g_s$?

Problem

Let q_s be family of Riemannian metrics on a manifold M with boundary, depending on a parameter $s \in (-\varepsilon, \varepsilon)$. If they all give the same spectrum, are there a diffeomorphisms $\phi_s : M \to M$ so that $q_0 = \phi_s^* q_s$?

In other words, are isospectral deformations necessarily trivial?

Problem

Let q_s be family of Riemannian metrics on a manifold M with boundary, depending on a parameter $s \in (-\varepsilon, \varepsilon)$. If they all give the same spectrum, are there a diffeomorphisms $\phi_s : M \to M$ so that $q_0 = \phi_s^* q_s$?

In other words, are isospectral deformations necessarily trivial?

This is within reach!

 Spectral rigidity has been previously proven on closed manifolds (compact, no boundary):

- Spectral rigidity has been previously proven on closed manifolds (compact, no boundary):
 - Negatively curved surfaces: Guillemin–Kazhdan 1980.
 - Negatively curved manifolds: Croke—Sharafutdinov 1998.
 - Anosov surfaces: Paternain—Salo—Uhlmann 2014.
 - Some more general manifolds: Paternain-Salo-Uhlmann 2015.

- Spectral rigidity has been previously proven on closed manifolds (compact, no boundary):
 - Negatively curved surfaces: Guillemin–Kazhdan 1980.
 - Negatively curved manifolds: Croke—Sharafutdinov 1998.
 - Anosov surfaces: Paternain-Salo-Uhlmann 2014.
 - Some more general manifolds: Paternain-Salo-Uhlmann 2015.
- We have adapted similar ideas of proof to manifolds with boundary.

Outline

- Seismic spectral data
- Spectra of a manifold with boundary
- 3 Different forms of uniqueness
- Spherical symmetry
 - Spherically symmetric manifolds
 - The Herglotz condition
- The main results
- 6 Anisotropy and geometry
- Ray transforms in low regularity

• Let $M=\bar{B}(0,1)\subset\mathbb{R}^n$ be the closed unit ball and c(x)=c(|x|) a $C^{1,1}$ sound speed.

- Let $M=\bar{B}(0,1)\subset\mathbb{R}^n$ be the closed unit ball and c(x)=c(|x|) a $C^{1,1}$ sound speed.
- \bullet The Riemannian metric on M is $g=c^{-2}(x)e.$ This makes (M,g) into a radially conformally Euclidean manifold.

- Let $M=\bar{B}(0,1)\subset\mathbb{R}^n$ be the closed unit ball and c(x)=c(|x|) a $C^{1,1}$ sound speed.
- The Riemannian metric on M is $g=c^{-2}(x)e$. This makes (M,g) into a radially conformally Euclidean manifold.
- If g is a rotation invariant Riemannian metric on M, there is a radial (more complicated if n=2) diffeomorphism $\phi\colon M\to M$ so that ϕ^*g is radially conformally Euclidean.

- Let $M=\bar{B}(0,1)\subset\mathbb{R}^n$ be the closed unit ball and c(x)=c(|x|) a $C^{1,1}$ sound speed.
- The Riemannian metric on M is $g=c^{-2}(x)e$. This makes (M,g) into a radially conformally Euclidean manifold.
- If g is a rotation invariant Riemannian metric on M, there is a radial (more complicated if n=2) diffeomorphism $\phi\colon M\to M$ so that ϕ^*g is radially conformally Euclidean.
- The Earth is spherically symmetric to a good approximation, but the best (elliptically anisotropic) radial model might not be conformally Euclidean. After a radial change of coordinates the metric becomes conformal — and Cartesian coordinates are lost.

Definition

A radial sound speed c(r) satisfies the Herglotz condition if

$$\frac{\mathrm{d}}{\mathrm{d}r} \left(\frac{r}{c(r)} \right) > 0$$

for all $r \in (0,1]$.

Definition

A radial sound speed c(r) satisfies the Herglotz condition if

$$\frac{\mathrm{d}}{\mathrm{d}r} \left(\frac{r}{c(r)} \right) > 0$$

for all $r \in (0, 1]$.

Equivalent formulations:

Definition

A radial sound speed c(r) satisfies the Herglotz condition if

$$\frac{\mathrm{d}}{\mathrm{d}r} \left(\frac{r}{c(r)} \right) > 0$$

for all $r \in (0, 1]$.

Equivalent formulations:

• All spheres $\{r = \text{constant}\}\$ are strictly convex. (Foliation condition!)

Definition

A radial sound speed c(r) satisfies the Herglotz condition if

$$\frac{\mathrm{d}}{\mathrm{d}r} \left(\frac{r}{c(r)} \right) > 0$$

for all $r \in (0,1]$.

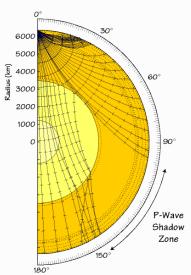
Equivalent formulations:

- All spheres $\{r = \text{constant}\}\$ are strictly convex. (Foliation condition!)
- The manifold is non-trapping and has strictly convex boundary.

• The radial Preliminary Reference Earth Model (PREM) is not $C^{1,1}$. Both pressure and shear wave speeds have jump discontinuities.

- The radial Preliminary Reference Earth Model (PREM) is not $C^{1,1}$. Both pressure and shear wave speeds have jump discontinuities.
- In addition, the shear wave speed vanishes in the liquid outer core.

- The radial Preliminary Reference Earth Model (PREM) is not $C^{1,1}$. Both pressure and shear wave speeds have jump discontinuities.
- In addition, the shear wave speed vanishes in the liquid outer core.
- Apart from these problems (jumps and liquid) both shear and pressure wave speeds do satisfy the Herglotz condition everywhere.



The Herglotz condition is satisfied: ray paths curve outwards. (Wikimedia Commons)

Outline

- Seismic spectral data
- Spectra of a manifold with boundary
- Oifferent forms of uniqueness
- Spherical symmetry
- The main results
 - Spectral rigidity
 - Length spectral rigidity
 - Ideas behind the proof
 - A numerical example of the trace formula
- 6 Anisotropy and geometry
- Ray transforms in low regularity

Theorem (de Hoop–I.–Katsnelson, 2017)

Theorem (de Hoop-I.-Katsnelson, 2017)

Let M be the closed unit ball in \mathbb{R}^3 . Let $c_s(r)$ be a family of radial sound speeds depending C^∞ -smoothly on both $s \in (-\varepsilon, \varepsilon)$ and $r \in [0, 1]$. Assume each c_s satisfies the Herglotz condition and a generic geometrical condition.

Theorem (de Hoop–I.–Katsnelson, 2017)

Let M be the closed unit ball in \mathbb{R}^3 . Let $c_s(r)$ be a family of radial sound speeds depending C^{∞} -smoothly on both $s \in (-\varepsilon, \varepsilon)$ and $r \in [0, 1]$. Assume each c_s satisfies the Herglotz condition and a generic geometrical condition.

If each c_s gives rise to the same spectrum (of the corresponding Laplace–Beltrami operator), then $c_s = c_0$ for all s.

Theorem (de Hoop–I.–Katsnelson, 2017)

Let M be the closed unit ball in \mathbb{R}^3 . Let $c_s(r)$ be a family of radial sound speeds depending C^{∞} -smoothly on both $s \in (-\varepsilon, \varepsilon)$ and $r \in [0, 1]$. Assume each c_s satisfies the Herglotz condition and a generic geometrical condition.

If each c_s gives rise to the same spectrum (of the corresponding Laplace–Beltrami operator), then $c_s = c_0$ for all s.

This simple model of the round Earth is spectrally rigid!

Corollary (de Hoop-I.-Katsnelson, 2017)

Corollary (de Hoop-I.-Katsnelson, 2017)

Let M be the closed unit ball in \mathbb{R}^3 . Let g_s be a family of rotation invariant metrics depending C^∞ -smoothly on $s \in (-\varepsilon, \varepsilon)$. Suppose each g_s is non-trapping with strictly convex boundary and assume a generic geometrical condition.

Corollary (de Hoop-I.-Katsnelson, 2017)

Let M be the closed unit ball in \mathbb{R}^3 . Let g_s be a family of rotation invariant metrics depending C^∞ -smoothly on $s \in (-\varepsilon, \varepsilon)$. Suppose each g_s is non-trapping with strictly convex boundary and assume a generic geometrical condition.

If the spectra of the Laplace–Beltrami operators Δ_{g_s} are all equal, then there is a family of radial diffeomorphisms $\phi_s\colon M\to M$ so that $\phi_s^*g_s=g_0$ for all s. That is, the manifolds (M,g_s) are isometric.

Theorem (de Hoop–I.–Katsnelson, 2017)

Theorem (de Hoop-I.-Katsnelson, 2017)

Let M be the closed unit ball in \mathbb{R}^n , $n\geq 2$. Let $c_s(r)$ be a family of radial sound speeds depending $C^{1,1}$ -smoothly on both $s\in (-\varepsilon,\varepsilon)$ and $r\in [0,1]$. Assume each c_s satisfies the Herglotz condition and a generic geometrical condition.

Theorem (de Hoop-I.-Katsnelson, 2017)

Let M be the closed unit ball in \mathbb{R}^n , $n\geq 2$. Let $c_s(r)$ be a family of radial sound speeds depending $C^{1,1}$ -smoothly on both $s\in (-\varepsilon,\varepsilon)$ and $r\in [0,1]$. Assume each c_s satisfies the Herglotz condition and a generic geometrical condition.

If each c_s gives rise to the same length spectrum, then $c_s = c_0$ for all s.

Theorem (de Hoop-I.-Katsnelson, 2017)

Let M be the closed unit ball in \mathbb{R}^n , $n\geq 2$. Let $c_s(r)$ be a family of radial sound speeds depending $C^{1,1}$ -smoothly on both $s\in (-\varepsilon,\varepsilon)$ and $r\in [0,1]$. Assume each c_s satisfies the Herglotz condition and a generic geometrical condition.

If each c_s gives rise to the same length spectrum, then $c_s = c_0$ for all s.

This simple model of the round Earth is length spectrally rigid!

Corollary (de Hoop-I.-Katsnelson, 2017)

Corollary (de Hoop-I.-Katsnelson, 2017)

Let M be the closed unit ball in \mathbb{R}^n , $n \geq 2$. Let g_s be a family of rotation invariant metrics depending $C^{1,1}$ -smoothly on $s \in (-\varepsilon, \varepsilon)$. Suppose each as is non-trapping with strictly convex boundary and satisfy a generic geometrical condition.

Corollary (de Hoop-I.-Katsnelson, 2017)

Let M be the closed unit ball in \mathbb{R}^n , $n \geq 2$. Let q_s be a family of rotation invariant metrics depending $C^{1,1}$ -smoothly on $s \in (-\varepsilon, \varepsilon)$. Suppose each as is non-trapping with strictly convex boundary and satisfy a generic geometrical condition.

If the length spectra of the manifolds (M, q_s) are all equal, then there is a family of radial (or more general if n=2) diffeomorphisms $\phi_s \colon M \to M$ so that $\phi_s^* q_s = q_0$ for all s. That is, the manifolds (M, q_s) are isometric.

Lemma (Trace formula)

Let $\lambda_0 < \lambda_1 \le \lambda_2 \le \dots$ be the positive eigenvalues of the Laplace–Beltrami operator. Define a function $f: \mathbb{R} \to \mathbb{R}$ by

$$f(t) = \sum_{k=0}^{\infty} \cos\left(\sqrt{\lambda_k} \cdot t\right).$$

Assume that the radial sound speed c satisfies some generic conditions.

Lemma (Trace formula)

Let $\lambda_0 < \lambda_1 \le \lambda_2 \le \dots$ be the positive eigenvalues of the Laplace–Beltrami operator. Define a function $f: \mathbb{R} \to \mathbb{R}$ by

$$f(t) = \sum_{k=0}^{\infty} \cos\left(\sqrt{\lambda_k} \cdot t\right).$$

Assume that the radial sound speed c satisfies some generic conditions.

The function f(t) is singular precisely at the length spectrum.

Lemma (Trace formula)

Let $\lambda_0 < \lambda_1 \le \lambda_2 \le \dots$ be the positive eigenvalues of the Laplace–Beltrami operator. Define a function $f: \mathbb{R} \to \mathbb{R}$ by

$$f(t) = \sum_{k=0}^{\infty} \cos\left(\sqrt{\lambda_k} \cdot t\right).$$

Assume that the radial sound speed c satisfies some generic conditions.

The function f(t) is singular precisely at the length spectrum.

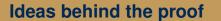
In particular, the spectrum determines the length spectrum.

Similar "trace formulas" and related results are known on closed manifolds (eg. Duistermaat–Guillemin 1975) and a weaker version on some manifolds with boundary (eg. Guillemin–Melrose 1979).

Similar "trace formulas" and related results are known on closed manifolds (eg. Duistermaat–Guillemin 1975) and a weaker version on some manifolds with boundary (eg. Guillemin–Melrose 1979).

Corollary

Spectral rigidity follows from length spectral rigidity.



Lemma

Let γ_s be a periodic broken ray (w.r.t. c_s) depending smoothly enough on s.

Lemma

Let γ_s be a periodic broken ray (w.r.t. c_s) depending smoothly enough on s. Then

$$\frac{\mathrm{d}}{\mathrm{d}s}\ell_s(\gamma_s) = \frac{1}{2} \int_{\gamma_s} \frac{\mathrm{d}}{\mathrm{d}s} c_s^{-2}.$$

Lemma

Let γ_s be a periodic broken ray (w.r.t. c_s) depending smoothly enough on s. Then

$$\frac{\mathrm{d}}{\mathrm{d}s}\ell_s(\gamma_s) = \frac{1}{2} \int_{\gamma_s} \frac{\mathrm{d}}{\mathrm{d}s} c_s^{-2}.$$

In particular, if the length spectrum does not depend on s, then $\frac{\mathrm{d}}{\mathrm{d}s}c_s^{-2}$ integrates to zero over (almost) all periodic broken rays.

Lemma (Periodic broken ray transform)

Assume the Herglotz condition. A radially symmetric function is uniquely determined by its integrals over (almost) all periodic broken rays.

Lemma (Periodic broken ray transform)

Assume the Herglotz condition. A radially symmetric function is uniquely determined by its integrals over (almost) all periodic broken rays.

Therefore $\frac{d}{ds}c_s^{-2}$ vanishes, and so c_s is independent of s.

Lemma (Periodic broken ray transform)

Assume the Herglotz condition. A radially symmetric function is uniquely determined by its integrals over (almost) all periodic broken rays.

Therefore $\frac{d}{ds}c_s^{-2}$ vanishes, and so c_s is independent of s.

This concludes the proof.

Lemma (Periodic broken ray transform)

Assume the Herglotz condition. A radially symmetric function is uniquely determined by its integrals over (almost) all periodic broken rays.

Therefore $\frac{d}{ds}c_s^{-2}$ vanishes, and so c_s is independent of s.

This concludes the proof.

Remark: No proof works without spherical symmetry.

A numerical example of the trace formula

Recall our proof:

A numerical example of the trace formula

Recall our proof:

 $\textcircled{ From eigenvalues } \lambda_0, \lambda_1, \lambda_2, \dots \text{ compute the function }$

$$f(t) = \sum_{k=0}^{\infty} \cos(\sqrt{\lambda_k} \cdot t).$$

This is the trace of the operator $\cos(\sqrt{-\Delta} \cdot t)$.

A numerical example of the trace formula

Recall our proof:

• From eigenvalues $\lambda_0, \lambda_1, \lambda_2, \ldots$ compute the function

$$f(t) = \sum_{k=0}^{\infty} \cos(\sqrt{\lambda_k} \cdot t).$$

This is the trace of the operator $\cos(\sqrt{-\Delta} \cdot t)$.

② See where f has singularities. The set of singularities is (more or less) the length spectrum.

Recall our proof:

• From eigenvalues $\lambda_0, \lambda_1, \lambda_2, \ldots$ compute the function

$$f(t) = \sum_{k=0}^{\infty} \cos(\sqrt{\lambda_k} \cdot t).$$

This is the trace of the operator $\cos(\sqrt{-\Delta} \cdot t)$.

- ② See where f has singularities. The set of singularities is (more or less) the length spectrum.
- Linearized length spectral data is periodic broken ray transform data.

Recall our proof:

• From eigenvalues $\lambda_0, \lambda_1, \lambda_2, \ldots$ compute the function

$$f(t) = \sum_{k=0}^{\infty} \cos(\sqrt{\lambda_k} \cdot t).$$

This is the trace of the operator $\cos(\sqrt{-\Delta} \cdot t)$.

- ② See where f has singularities. The set of singularities is (more or less) the length spectrum.
- Linearized length spectral data is periodic broken ray transform data.
- The periodic broken ray transform can be inverted explicitly for radial functions.

Example:

Example: We want to find the length of an interval, given its Neumann spectrum.

Example: We want to find the length of an interval, given its Neumann spectrum.

If the length of the interval is L, the eigenvalues are

$$\lambda_k = \left(\frac{k\pi}{L}\right)^2, \quad k = 0, 1, 2, \dots$$

Suppose $L=\frac{1}{2}$ and we have measured the numbers $0,4\pi^2,16\pi^2,\dots$

Example: We want to find the length of an interval, given its Neumann spectrum.

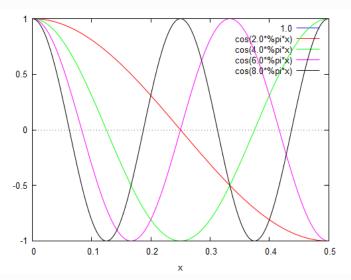
If the length of the interval is L, the eigenvalues are

$$\lambda_k = \left(\frac{k\pi}{L}\right)^2, \quad k = 0, 1, 2, \dots$$

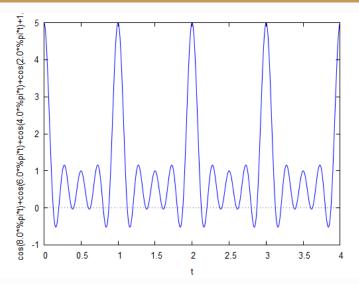
Suppose $L=\frac{1}{2}$ and we have measured the numbers $0,4\pi^2,16\pi^2,\dots$

We compute and plot the trace function

$$f(t) = \sum_{k=0}^{\infty} \cos(\sqrt{\lambda_k} \cdot t).$$



Eigenfunctions for k = 0, 1, 2, 3, 4.



Trace function computed from k = 0, 1, 2, 3, 4.

Outline

- Seismic spectral data
- Spectra of a manifold with boundary
- Oifferent forms of uniqueness
- Spherical symmetry
- The main results
- 6 Anisotropy and geometry
 - Elliptic and general elastic anisotropy
 - Pressure and shear waves
 - Anisotropy and coordinates
 - Our model
- Ray transforms in low regularity

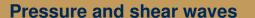
 A material is anisotropic if sound speed depends on direction. There are different types of direction dependence:

- A material is anisotropic if sound speed depends on direction. There are different types of direction dependence:
 - General elliptic anisotropy corresponds to a Riemannian manifold (a manifold with a Riemann metric).

- A material is anisotropic if sound speed depends on direction. There are different types of direction dependence:
 - General elliptic anisotropy corresponds to a Riemannian manifold (a manifold with a Riemann metric).
 - General anisotropy corresponds to a Finsler manifold (a manifold with a Finsler metric).

- A material is anisotropic if sound speed depends on direction. There are different types of direction dependence:
 - General elliptic anisotropy corresponds to a Riemannian manifold (a manifold with a Riemann metric).
 - General anisotropy corresponds to a Finsler manifold (a manifold with a Finsler metric).
 - Riemannian manifolds are a very special subclass of Finsler manifolds.

- A material is anisotropic if sound speed depends on direction. There are different types of direction dependence:
 - General elliptic anisotropy corresponds to a Riemannian manifold (a manifold with a Riemann metric).
 General anisotropy corresponds to a Finsler manifold (a manifold with
 - General anisotropy corresponds to a Finsler manifold (a manifold with a Finsler metric).
 - Riemannian manifolds are a very special subclass of Finsler manifolds.
- A material is isotropic if sound speed is independent of direction. This can be modeled by a conformally Euclidean metric.



Pressure and shear waves

 There are pressure and shear waves in an elastic medium, and they have different sound speeds.

Pressure and shear waves

- There are pressure and shear waves in an elastic medium, and they have different sound speeds.
- To model elastic waves in general anisotropy, one needs a manifold with two Finsler metrics, one for pressure and one for shear waves.

Pressure and shear waves

- There are pressure and shear waves in an elastic medium, and they have different sound speeds.
- To model elastic waves in general anisotropy, one needs a manifold with two Finsler metrics, one for pressure and one for shear waves.
- In fact, the shear wave speed might not even by a Finsler metric in the traditional sense.

• Let $\phi \colon M \to M$ be a diffeomorphism of a manifold that keeps the boundary fixed.

- Let $\phi \colon M \to M$ be a diffeomorphism of a manifold that keeps the boundary fixed.
- If g (or F) is a Riemannian (or Finsler) metric on M, then the pullback ϕ^*g (or ϕ^*F) is different Riemannian metric that behaves exactly the same for boundary measurements.

- Let $\phi \colon M \to M$ be a diffeomorphism of a manifold that keeps the boundary fixed.
- If g (or F) is a Riemannian (or Finsler) metric on M, then the pullback ϕ^*g (or ϕ^*F) is different Riemannian metric that behaves exactly the same for boundary measurements.
- A fully anisotropic model can never be reconstructed from boundary measurements uniquely. The data is always invariant under changes of coordinates.

- Let $\phi \colon M \to M$ be a diffeomorphism of a manifold that keeps the boundary fixed.
- If g (or F) is a Riemannian (or Finsler) metric on M, then the pullback ϕ^*g (or ϕ^*F) is different Riemannian metric that behaves exactly the same for boundary measurements.
- A fully anisotropic model can never be reconstructed from boundary measurements uniquely. The data is always invariant under changes of coordinates.
- The best one can hope for is reconstruction up to changes of coordinates.

No S-waves. — Only one metric.

- No S-waves. Only one metric.
- Isotropic P-wave speed. Conformally Euclidean metric.

- No S-waves. Only one metric.
- Isotropic P-wave speed. Conformally Euclidean metric.
- Spherical symmetry.

- No S-waves. Only one metric.
- Isotropic P-wave speed. Conformally Euclidean metric.
- Spherical symmetry.
- Reconstruction possible in the natural Cartesian coordinates. No gauge freedom.

Outline

- Seismic spectral data
- Spectra of a manifold with boundary
- 3 Different forms of uniqueness
- Spherical symmetry
- The main results
- 6 Anisotropy and geometry
- Ray transforms in low regularity
 - X-ray transforms
 - Periodic broken ray transforms

X-ray transforms

X-ray transforms

Theorem (de Hoop-I., 2017)

Let M be a rotation symmetric non-trapping manifold with a piecewise $C^{1,1}$ metric and strictly convex boundary. Then the geodesic X-ray transform is injective on $L^2(M)$.

X-ray transforms

Theorem (de Hoop-I., 2017)

Let M be a rotation symmetric non-trapping manifold with a piecewise $C^{1,1}$ metric and strictly convex boundary. Then the geodesic X-ray transform is injective on $L^2(M)$.

Earlier similar results:

- The X-ray transform (Radon et al.): Euclidean metric (c is constant).
- Mukhometov, 1977: Smooth simple metrics (simplicity is stronger than Herglotz).
- Sharafutdinov, 1997: C^{∞} metrics and C^{∞} functions.

Periodic broken ray transforms

Periodic broken ray transforms

Theorem (de Hoop-I., 2017)

Let M be a rotation symmetric non-trapping manifold with a $C^{1,1}$ metric and strictly convex boundary and dimension at least three. Assume that there are not too many conjugate points at the boundary. The integrals of a function $f \in L^p(M)$, p > 3, over all periodic broken rays determines the even part of the function.

Very little can be recovered of the odd part.

Periodic broken ray transforms

Theorem (de Hoop-I., 2017)

Let M be a rotation symmetric non-trapping manifold with a $C^{1,1}$ metric and strictly convex boundary and dimension at least three. Assume that there are not too many conjugate points at the boundary. The integrals of a function $f \in L^p(M)$, p > 3, over all periodic broken rays determines the even part of the function.

Very little can be recovered of the odd part.

Tools used:

- Planar average ray transform.
- Abel transform.
- Funk transform.
- Fourier series.

DISCOVERING MATH at JYU. Since 1863.

Slides and papers available at http://users.jyu.fi/~jojapeil