

Spectral and travel time tomography on Mars with InSight

Applied Inverse Problems minisymposium

How to see inside the Earth? Theory and applications of seismic inverse problems

Joonas Ilmavirta

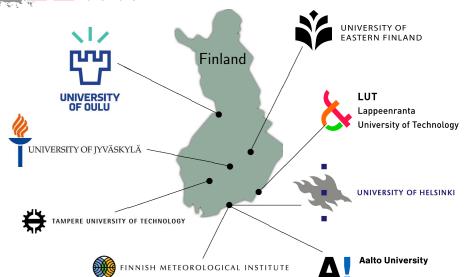
July 12, 2019

Based on joint work with

Maarten de Hoop and Vitaly Katsnelson







Conference announcement

The annual Finnish inverse problems conference "Inverse Days" will be organized in Jyväskylä 16–18 December, 2019.

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http://r.jyu.fi/yVK
(https://www.jyu.fi/science/en/maths/research/inverse-problems/id2019/)
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Registration opens this week!

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 What can you say for sure and is there an inversion algorithm?
- Identifying useful data sets can help future mission planning.
- Grand goal: A mathematical theory of seismic planetary exploration.

Outline

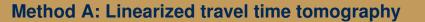
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- Seeing an entire planet

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- Mars is roughly spherically symmetric. There are reliable ways to reconstruct a radial model of the (upper) mantle from a single station. (The mantle determines the CMB.)
- I will ignore noise, model errors, finiteness, stability, and many other practical things.



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- Data: Pairs of directions (≈ angle from normal) and times.
 Uknown: Wave speed (≈ geometry).
- The set of all periodic travel times is the length spectrum.

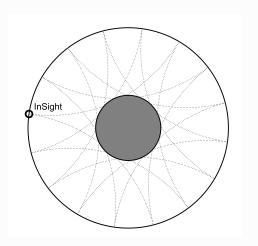
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- Solution: Linearize!
- Linearized data: Pairs of periodic broken rays and integrals over them.
 Uknown: Variations of wave speed (a function).



Periodic seismic ray reflecting on the surface and CMB.

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(Uniqueness should be provable for the non-linear one, too.)

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- The spectrum of free oscillations can be measured from any single point.
- Mathematically, the spectrum of free oscillations corresponds to the Neumann spectrum of the Laplace—Beltrami operator on a manifold.

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If a family of wave speeds $c_s(r)$ have the same spectrum, are the equal? Is the (Martian) mantle spectrally rigid?

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This is based on linearization. With a trace formula one ends up showing that the Martian mantle is length spectrally rigid.

 Seismic events with known sources are another source of information, and the most useful type seems to be meteorite impacts.

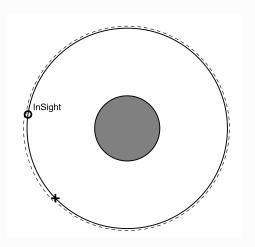
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- If there are no other events on the same great circle around the same time, we can measure the time difference δ .
- Multiple arrivals or a priori information tells the time T around the great circle.



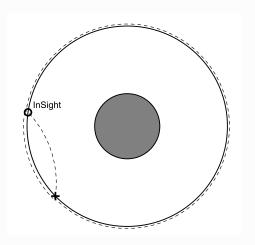
Two surface wave arrivals from the same event.

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- Assuming the seismometer can detect directions of surface wave arrivals, we can deduce the time and place of the event.
- This was all done on surface, and it gives rise to interior data: Now using body waves we know the travel time between InSight and the source.
- To get here, we needed to assume spherical symmetry only on the surface, but the arising problem is easiest to solve if the symmetry extends inside.



The body wave whose initial point and time were located with surface waves.

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- The linearized problem is X-ray tomography (or an Abel transform), and can also be solved explicitly. (e.g. de Hoop-I., 2017)

• We have three methods to obtain the wave speed c(r) in the mantle down to the depth where the Herglotz condition first fails.

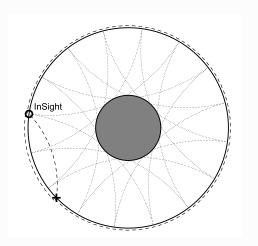
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- This gives us an isotropic radially symmetric reference model of the mantle, which is a stepping stone towards deeper and finer structure.



Three ways to see the mantle from InSight.

- A: From noise correlations to (linearized) travel times.
- B: From spectrum to length spectrum.
- C: Meteorites; body wave data calibrated by surface waves.

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- Seeing an entire planet

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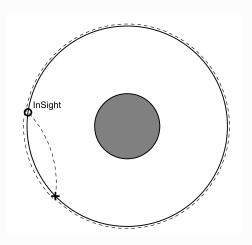
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- It remains to be proven that this is indeed the only obstruction to uniqueness.

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- We assumed that the surface is spherically symmetric (or otherwise known), but we needed no assumption on the interior.
- This leads to travel time data: The travel times (geometrically: distances) are known from all points on the surface to a single fixed point.



The body wave whose initial point and time were located with surface waves.

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This is possible at least in Euclidean geometry or with real analytic perturbations but always unstable.

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- Most geometrical inverse problems work with smooth manifolds. How to add conormal singularities and finite interior regularity?
- How does spectral rigidity and X-ray tomography work in a rough onion?
- What is the geometry of periodic broken rays?

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- We do not have

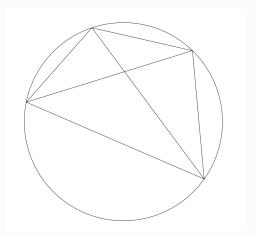
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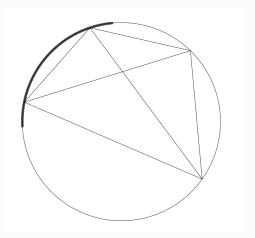
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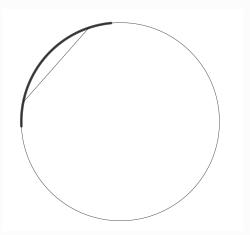


Boundary distance rigidity: Do the distances between all boundary points determine the meometry?

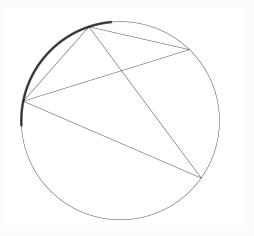


We have an accessible region — a measurement array.

The size is exaggerated.



In the local boundary distance problem one knows the distances between the points in the small set and wants to find the geometry near that set.



The "half-local" boundary distance data has more information and one wants to reconstruct the whole geometry.

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