Geometric inverse problems arising from geophysics Seminar at Tampere

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Positions

• Aug 2020-: Assistant professor at Tampere

- Aug 2020-: Assistant professor at Tampere
- Sep 2020-: Academy research fellow at Tampere

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Geometric inverse problems arising from geophysics



Outline

Mathematical modelling

- Inverse modelling and imaging
- Elastic geometry
- Imaging with neutrinos
- Geometrization of gravitation
- Elastic waves
- Elastic geometry
- 5 Inverse problems

Inverse modelling and imaging

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I will give two examples.

Elastic geometry

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Enough generality to allow for known physical phenomena and maybe a little more.

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- are slow enough and fast enough,
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- actually exist, and
- can be detected.

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I am studying geometric neutrino data with Gunther Uhlmann.

Outline

Mathematical modelling

- 2 Geometrization of gravitation
 - Newton's theory
 - Einstein's theory
 - The goal

Elastic waves

- Elastic geometry
- 5 Inverse problems

Newton's theory

Joonas Ilmavirta (Tampere)

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- The gravitational force exerted by the Sun causes the Earth's trajectory to curve.
- The force is described by a simple formula and the equation of motion is an ODE in ℝⁿ.
- The Newtonian approach is straightforward to use and often a good model.

Einstein's theory

Joonas Ilmavirta (Tampere)

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- There is a complicated equation of motion for the geometry itself: Einstein's field equation is a non-linear system of coupled PDEs.
- This model is harder to use but can reach phenomena inaccessible to Newtonian gravity and provides a more geometric way to see the essential structures.

A geometric theory of elasticity?

Outline

- Mathematical modelling
- Geometrization of gravitation

Elastic waves

- The stiffness tensor
- The elastic wave equation
- The principal symbol
- Polarization
- Singularities and the slowness surface

Elastic geometry

Inverse problems

The stiffness tensor

Joonas Ilmavirta (Tampere)

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- Density normalized: $a_{ijkl}(x) = c_{ijkl}(x)/\rho(x)$.

• Using Newton's second law with a restoring force given by Hooke's law leads to the elastic wave equation (EWE)

$$\sum_{j,k,l} \partial_j [c_{ijkl}(x)\partial_k u_l(x,t)] - \rho(x)\partial_t^2 u_i(x,t) = 0,$$

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- If the material is anisotropic (*c* is no more symmetric than necessary and wave speed depends on direction), then the vector nature of the equation cannot be ignored.
- Elastic waves arising from earthquakes (or marsquakes!) satisfy this equation away from the focus of the event to great accuracy.

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• The principal symbol of the EWO is $\Gamma(x,\xi) - \omega^2 I$, where $\xi = \omega p$.

Polarization

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- In anisotropic elasticity it does not work quite as nicely. The fastest polarization is called quasi-P and the slower ones quasi-S.
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- Decomposition to polarizations only works on the level of singularities. The individual polarizations do not satisfy PDEs.

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• The admissible slowness vectors *p* are on the slowness surface given by the equation

$$\det(\Gamma(x,p) - I) = 0.$$



The slowness surface. Smaller slowness \iff faster wave.

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- Elastic waves
- Elastic geometry
 - Distance
 - Ray tracing
 - Finsler manifolds

Inverse problems

Distance

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- That is, we declare the distance between x and y to be the shortest amount of time it takes for a wave to travel from x to y. (qP is fastest!)

Ray tracing

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- We can then study how these particles travel.
 - Traditional view: The trajectory of the phonon is curved because wave speed varies.
 - Newer view: The phonon goes straight in a curved geometry (geodesic), and the geometry is curved by variations in wave speed.

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- In elastic geometry we measure distance in travel time, and the waves go straight in this geometry.
- Fermat's principle: Phonons the particles corresponding to elastic waves — go straight in the geometry given by travel time.
- Fermat's principle is about going straight in the relevant geometry, not about taking the shortest path. These are not the same thing over long distances or for shear waves.

Finsler manifolds

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- The norms on the dual spaces T_x^*M have the same properties.

Elastic finsler geometry

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Other polarizations are problematic.

Outline

- Mathematical modelling
- Geometrization of gravitation
- Elastic waves
- Elastic geometry
- Inverse problems
 - Geometrization of the question
 - Herglotz (Mönkkönen)
 - Dix (de Hoop, Lassas)
 - Distance function (de Hoop, Lassas, Saksala)
 - Scattering data (de Hoop, Lassas, Saksala)
 - Ray tracing (Iversen, Ursin, Saksala, de Hoop)
 - And more...

Geometrization of the question

• Typical inverse problem: Given some boundary data, find the reduced stiffness tensor $\rho^{-1}(x)c_{ijkl}(x)$ everywhere.

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- From the slowness surface one can then find the material parameters

 the components of the stiffness tensor.
- Practical goal for a geometer: Given some boundary data, find the Finsler manifold or its cosphere bundle.

Herglotz (Mönkkönen)

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- Linearized travel time data leads to X-ray tomography.

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- The "directionality" of Finsler geometry is a major complication in comparison to the Riemannian version (de Hoop–Holman– Iversen–Lassas–Ursin, 2015).

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- If *F* is fiberwise real analytic (elasticity or Riemann!), then *F* is determined uniquely.

Scattering data (de Hoop, Lassas, Saksala)

Joonas Ilmavirta (Tampere)

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- This broken scattering relation can see much more of *TM*, but the trapped set is still invisible.
- Global uniqueness is can be done with added assumptions: reversibility (point symmetry) and foliation.
- Almost no assumptions are needed in the Riemannian case (Kurylev–Lassas–Uhlmann, 2010).

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 - Fermi coordinates and covariant derivatives of the corresponding elastic geometry are tricky to set up but the equations are clean.
- Variations in position (Q) and momentum (P) satisfy an equation

$$\partial_t \begin{pmatrix} Q(t) \\ P(t) \end{pmatrix} = \begin{pmatrix} W^T(t) & V(t) \\ -U(t) & -W(t) \end{pmatrix} \begin{pmatrix} Q(t) \\ P(t) \end{pmatrix}.$$

- We follow seismic rays and study their variations.
- There are different coordinates:
 - Cartesian coordinates where things are trivial to define but equations are messy.
 - Ray-centered coordinates which are more complicated to use but more structure arises.
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- Variations in position (Q) and momentum (P) satisfy an equation

$$\partial_t \begin{pmatrix} Q(t) \\ P(t) \end{pmatrix} = \begin{pmatrix} W^T(t) & V(t) \\ -U(t) & -W(t) \end{pmatrix} \begin{pmatrix} Q(t) \\ P(t) \end{pmatrix}.$$

• Written in terms of a Jacobi field *J* and its covariant derivative, we have instead

$$D_t \begin{pmatrix} J(t) \\ D_t J(t) \end{pmatrix} = \begin{pmatrix} 0 & I \\ -R(t) & 0 \end{pmatrix} \begin{pmatrix} J(t) \\ D_t J(t) \end{pmatrix}$$

And more...

Joonas Ilmavirta (Tampere)

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Modelling goal

Building a complete theory of elastic geometry.

Key ideas:

- Pure mathematics for the sake of physics.
- Geometrization.
- Geomathematics.

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